



Ab Initio Phase Diagrams of Minerals

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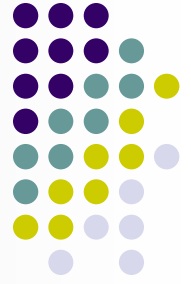


Daniel
Y.
Jung

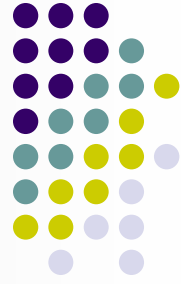


Donat
Adams

Acknowledgements:



- Collaborators: S.Ono, M.Gillan, D.Price, ABINIT group, VASP group, ...
- Supercomputers: CSCS (Manno), CSAR (Manchester), own.
- Synchrotron: SPring8 (Japan).
- ETH Zurich: funding.



Plan:

- Introduction
- Simulation Methods
- Results

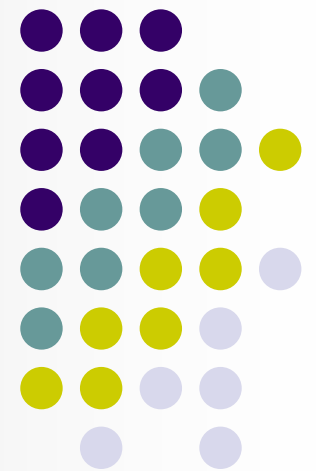
MgO

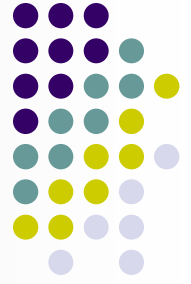
SiO₂

MgSiO₃ perovskite and post-perovskite

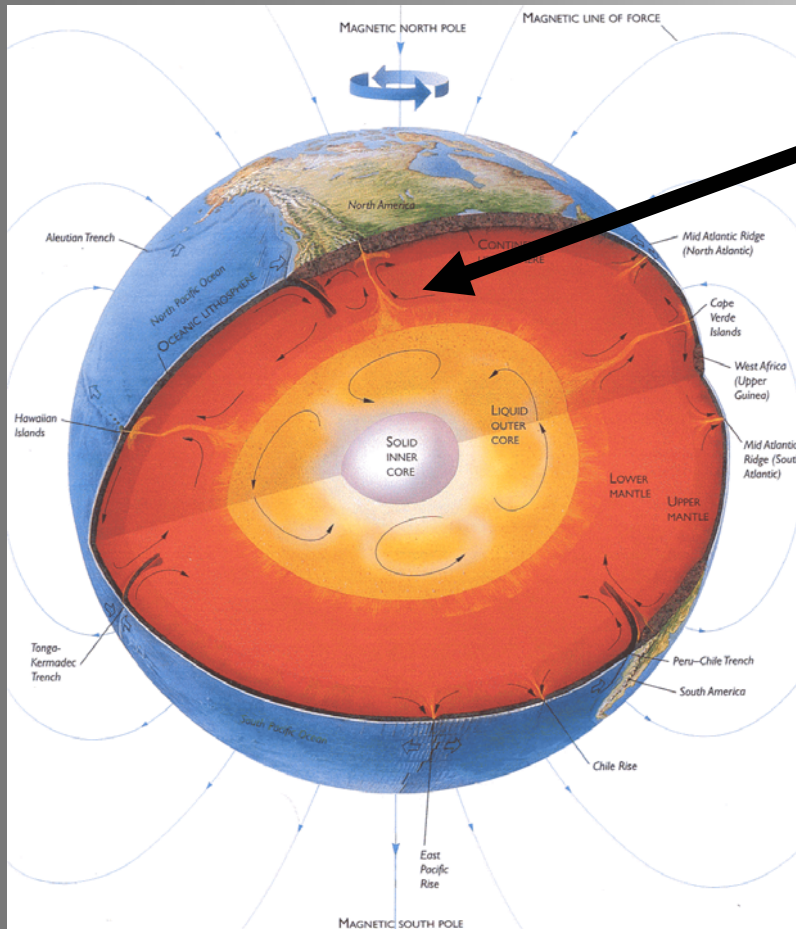
- Conclusions
- Appendices

1. Introduction





Earth's Interior

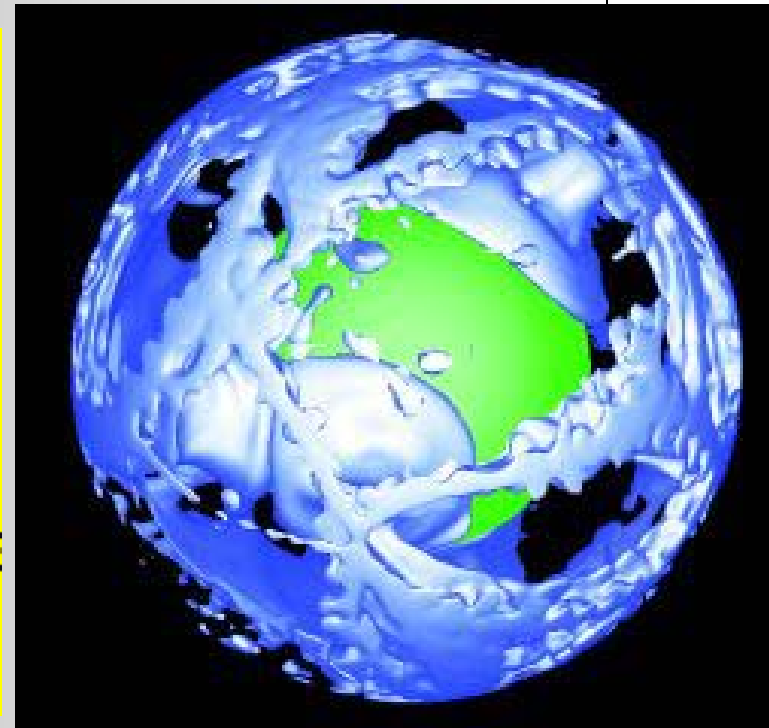
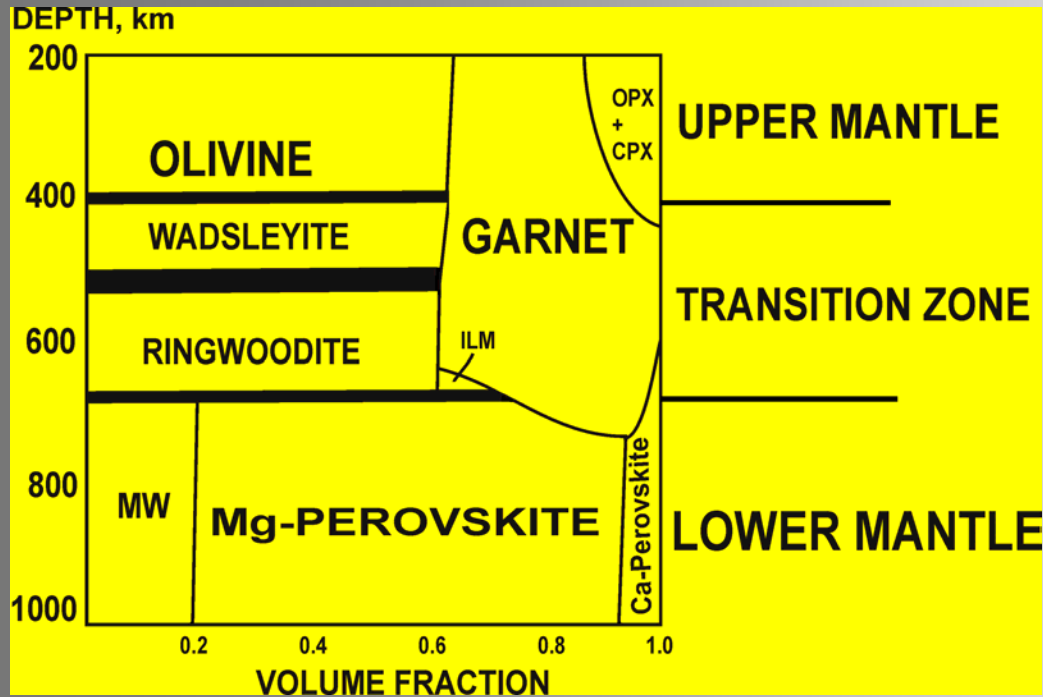


Earth's structure

Lower mantle:

- 53 vol.% of the Earth!!!
- 75 vol.% - MgSiO_3 perovskite
- 20% - $(\text{Mg,Fe})\text{O}$
- 5% - CaSiO_3 perovskite
- Bottom 200 km -**MYSTERY!!!** (D'' layer).

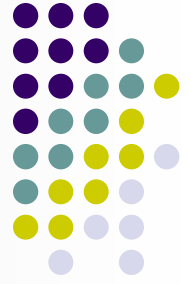
Phase Diagrams and Deep Earth



Phase relations in mantle
(after Ringwood, 1991)

Mantle convection: 670 km boundary
is a partial barrier
(thanks to P.Tackley)

Seismic discontinuities. Geochemistry. Geodynamics.

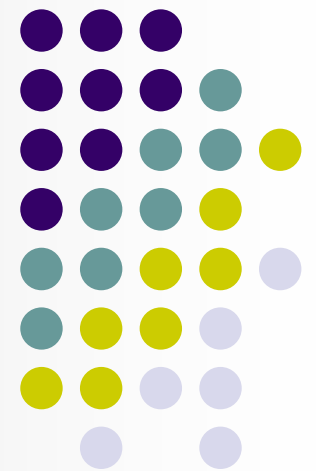


Which Diagrams Needed?

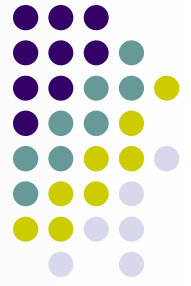
- Mantle: $\text{MgO-SiO}_2\text{-CaO-Al}_2\text{O}_3\text{-FeO-Fe}_2\text{O}_3$.
- p, T, x -phase diagrams, ideally.

Here: pure MgO , SiO_2 , MgSiO_3 .

2. Simulation Methods



Density-Functional Perturbation Theory

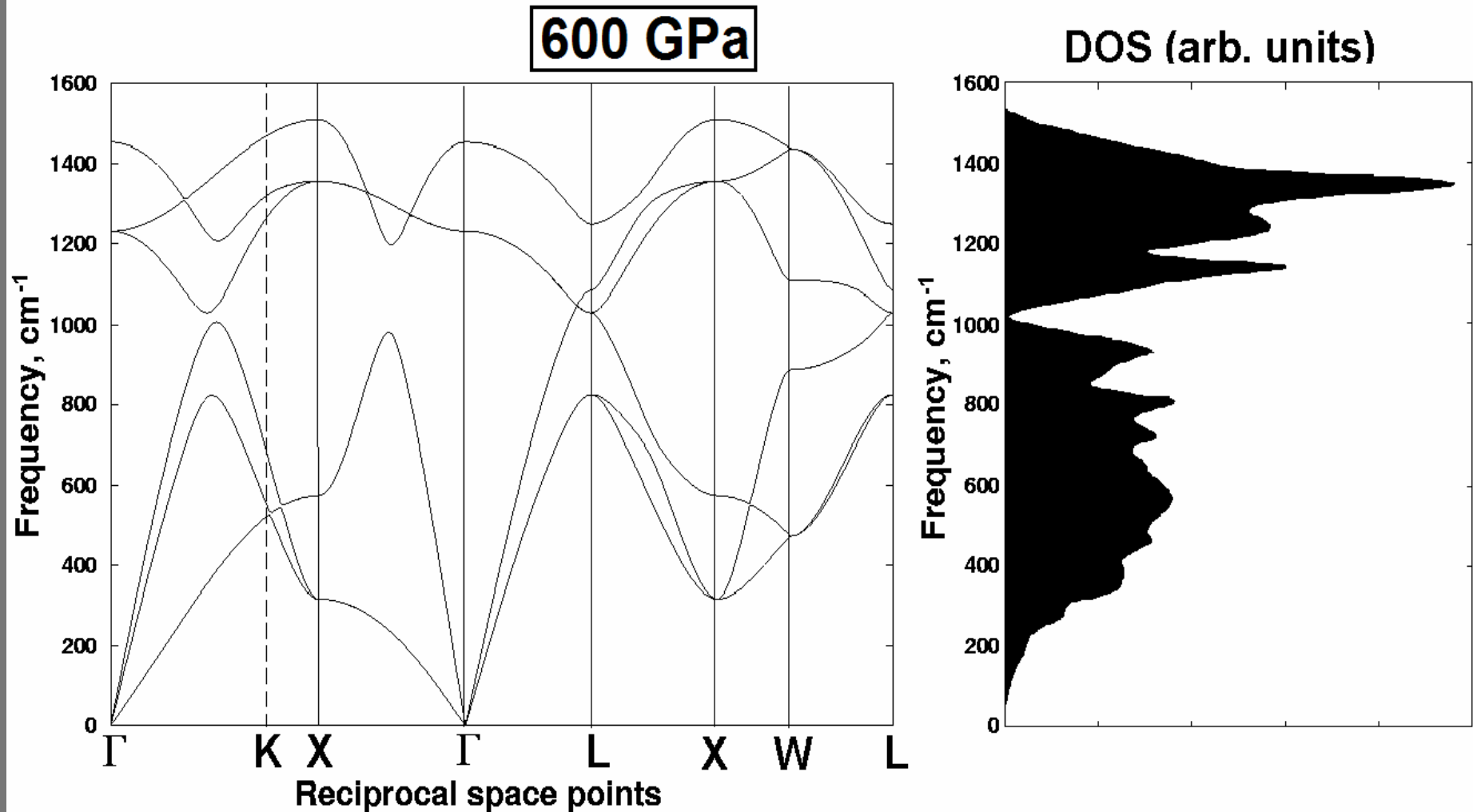


- Quasiharmonic approximation.
- Phonon frequencies $\omega_i(\mathbf{k})$ and density of states $g(\omega)$.
- Thermodynamic properties, e.g.:

$$F(T) = E_0 + \int_0^{\omega_{\max}} \frac{1}{2} \hbar \omega g(\omega) d\omega + k_B T \int_0^{\omega_{\max}} \ln[1 - \exp(-\frac{\hbar \omega}{k_B T})] g(\omega) d\omega$$

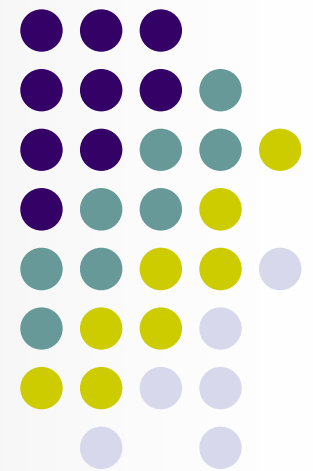
- $G = F + (-dF/dV)V$.
- ABINIT code.

Ab Initio Lattice Dynamics



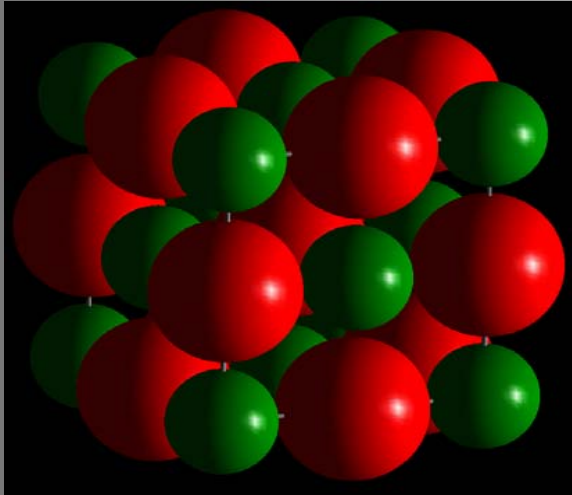
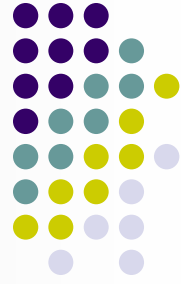
Phonon dispersion curves and phonon density of states of MgO: 0 GPa, 400 GPa, 600 GPa. (Oganov et al., J.Chem.Phys. **118**, 10174 (2003))

3. Results

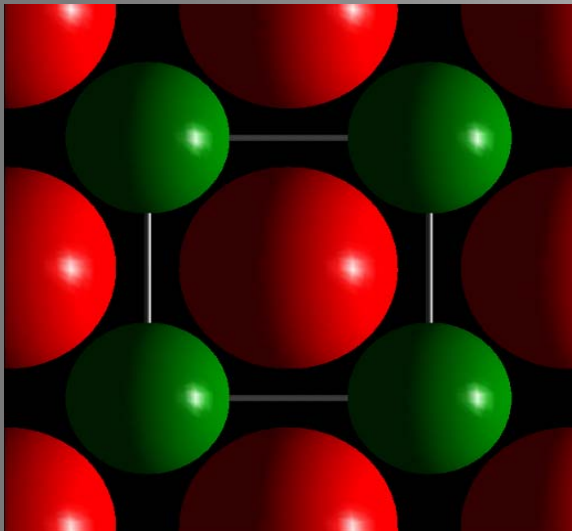


Story 1: MgO

B1 vs B2 (not Vitamins!)



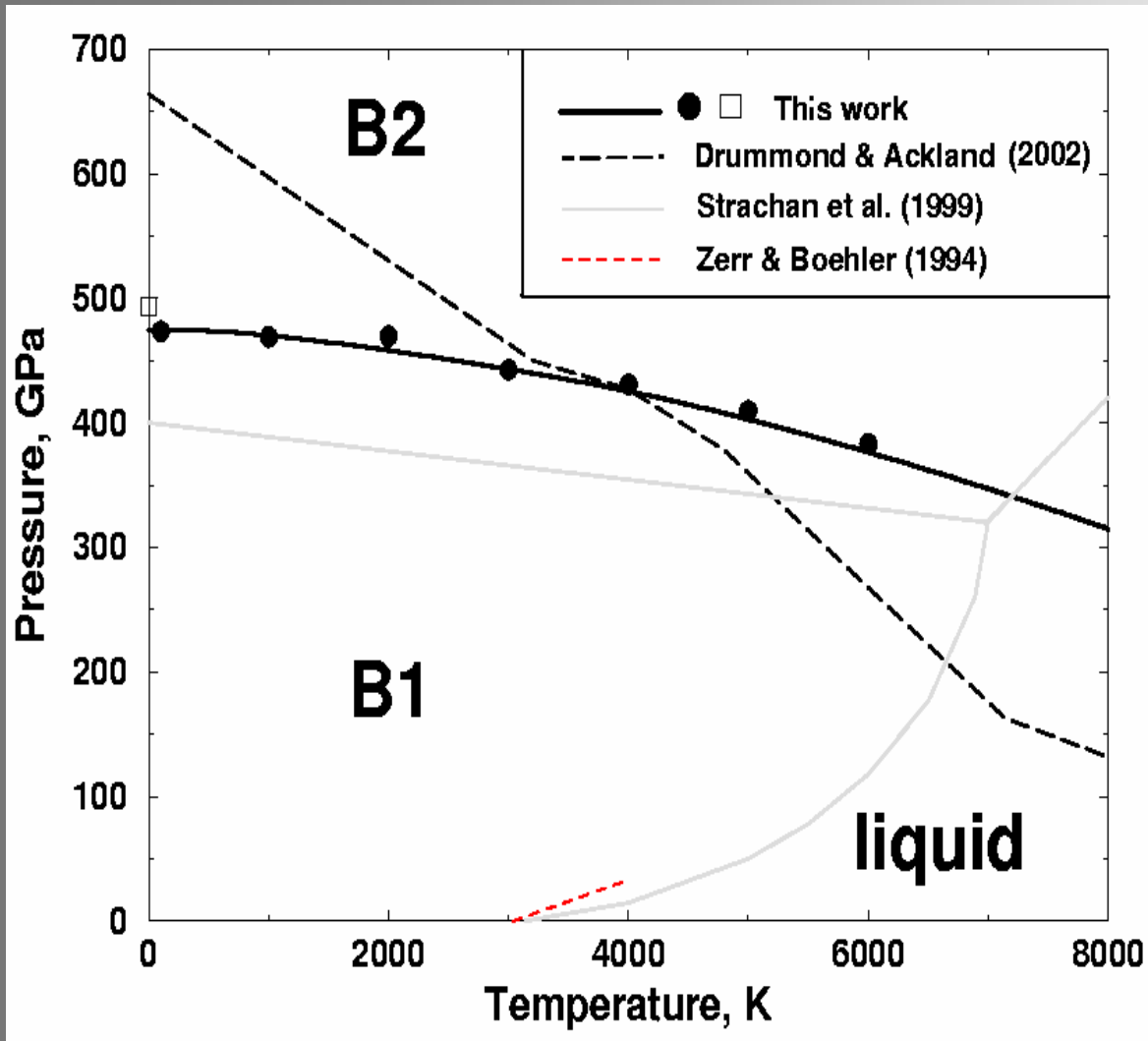
B1



B2

- B1 – structure type NaCl.
- B2 – structure type CsCl.
- Unique stability of the B1 structure!
- Ideal as a pressure calibrant.

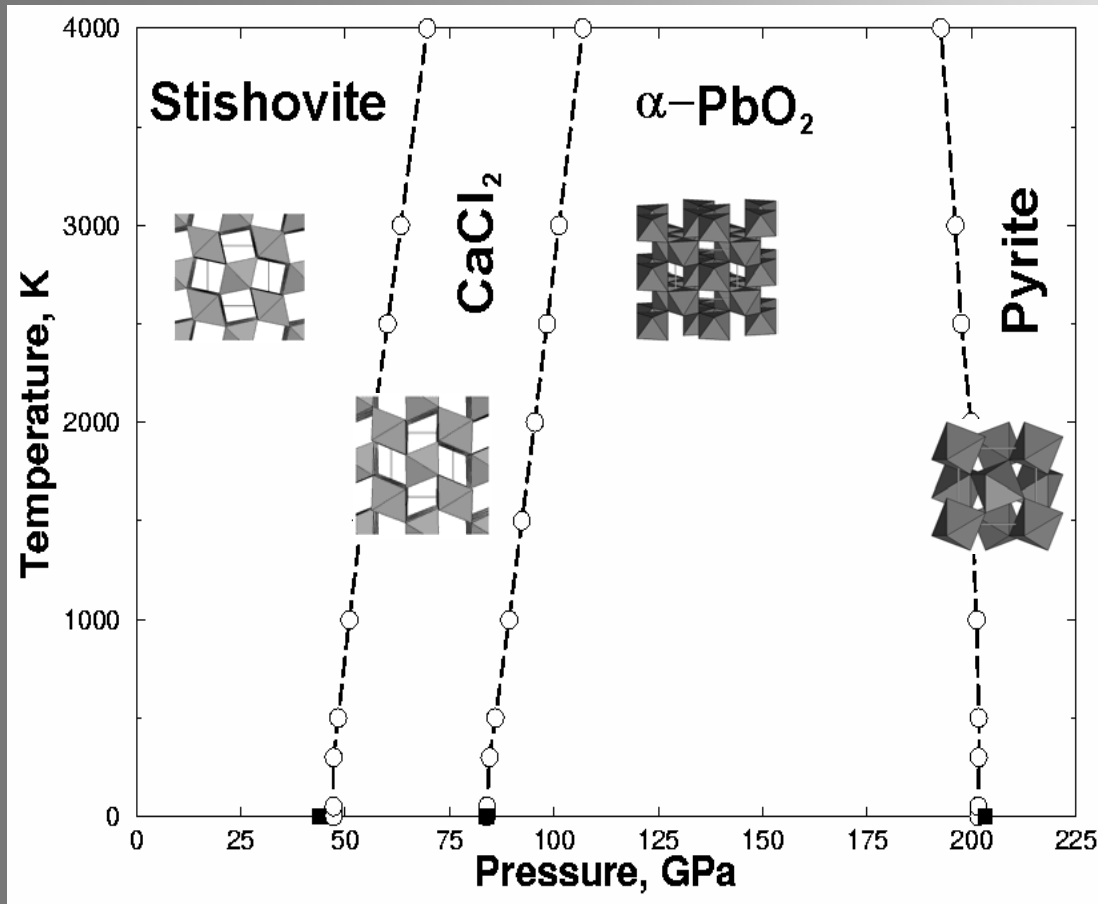
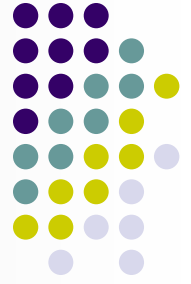
Phase diagram of MgO



◆ Only B1 phase in the Earth.

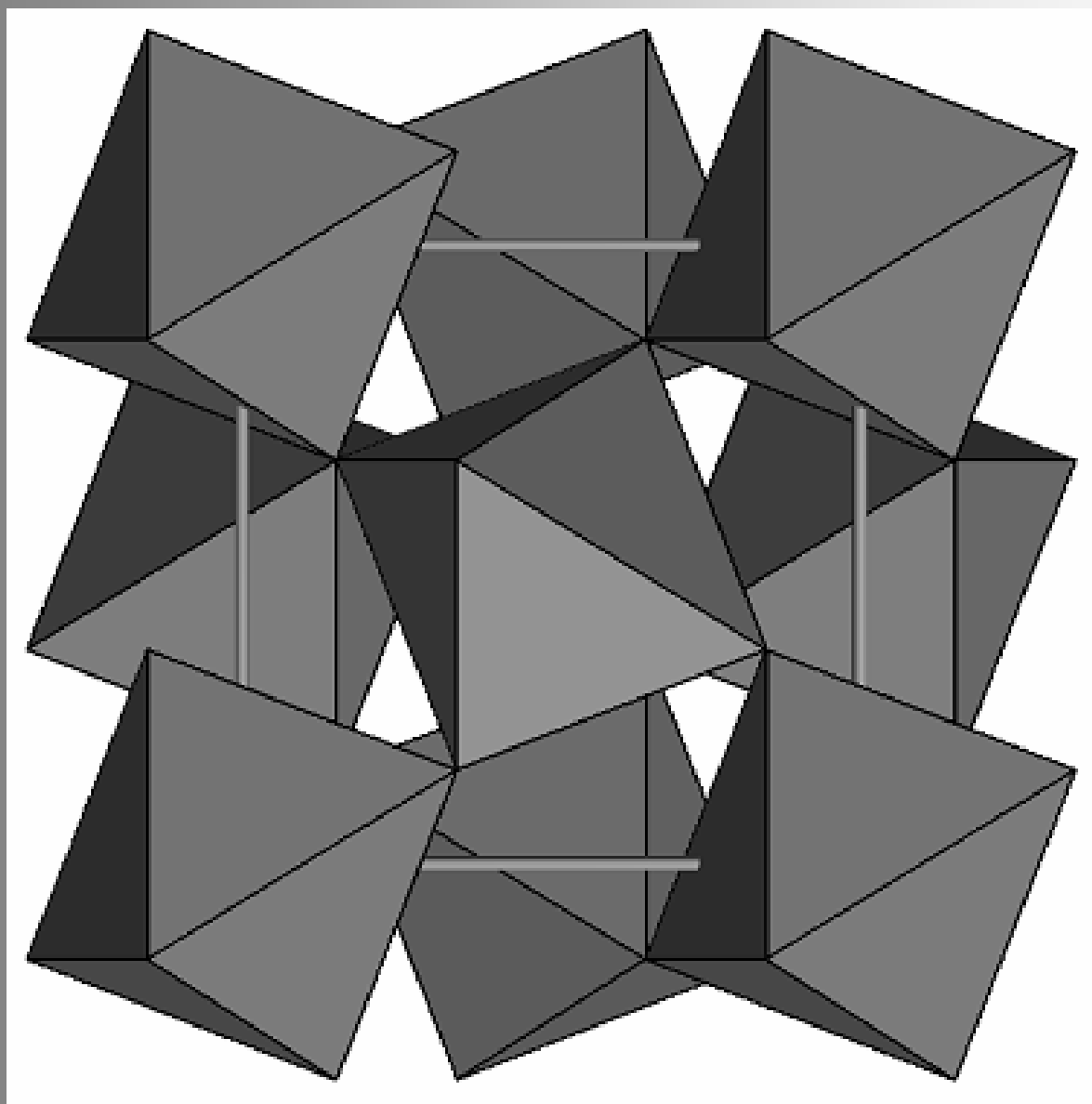
Oganov et al., J.Chem.Phys. **118**, 10174 (2003)

Story 2: Phase diagram of SiO_2



Oganov et al., Submitted to PRL (2004)

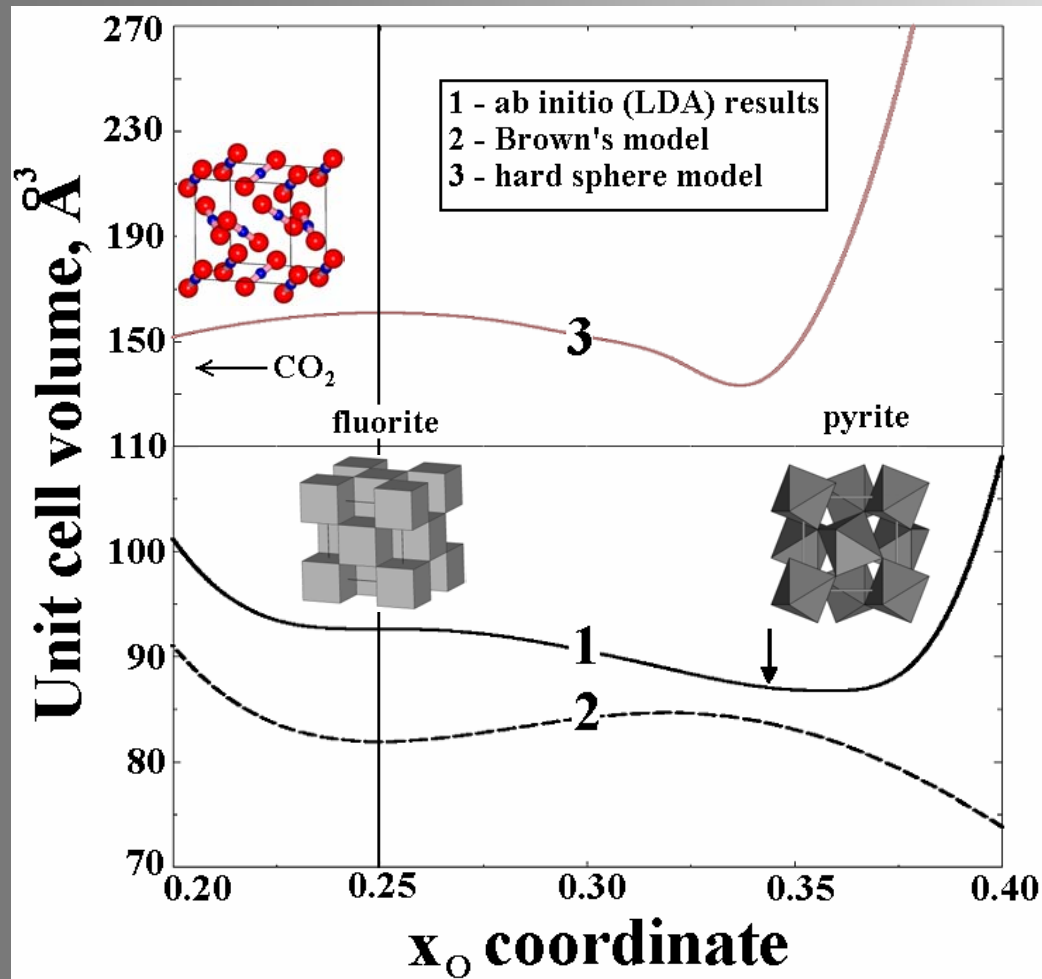
- No seismic discontinuities from SiO_2 transitions.
- Breakdown of close packing at >200 GPa!



Pyrite-type structure ($210 \text{ GPa} < P$)



Breakdown of Close Packing



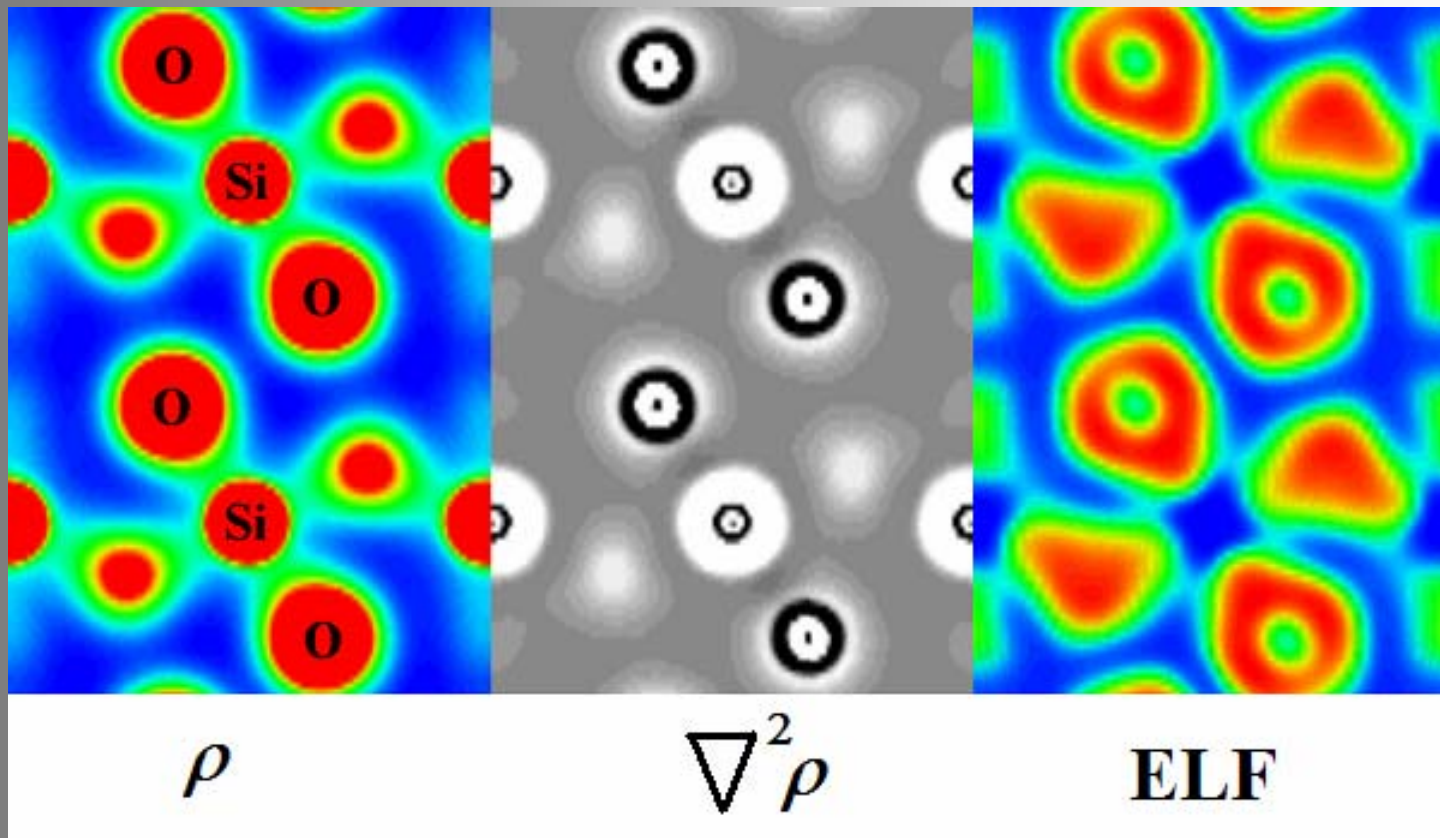
From Oganov et al., submitted to PRL (2004).

- Atoms don't behave as spheres!
- Non-close-packed structures far denser!

Seeing Atoms and Bonds

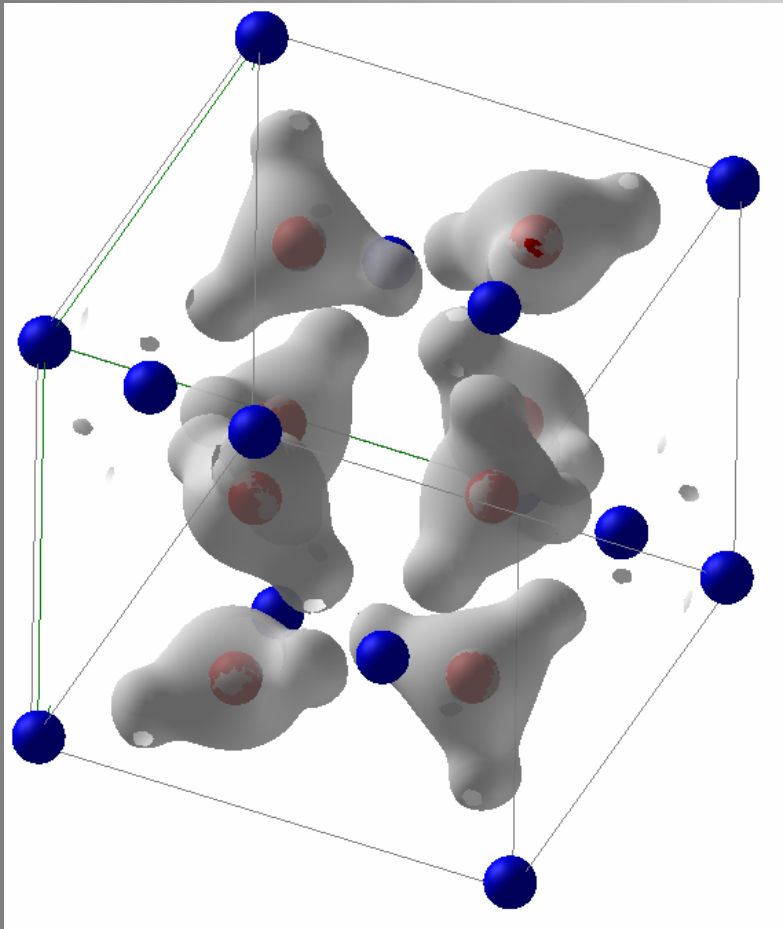
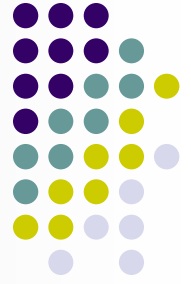


- Aspherical atoms.
- No O-O bonds, despite (3,-1) critical points.
- Si charge: +3.17 (Bader), +4.02 (Born).



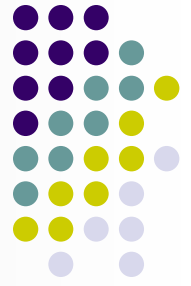
Oganov et al., submitted to PRL (2004).

Localised Orbital Locator



- Schmider & Becke (2000,2002).
- Own implementation in VASP.

Story 3: MgSiO_3 perovskite



Perovskite crystals (CaTiO_3)

- ◆ $(\text{Mg,Fe})\text{SiO}_3$ perovskite ~40 vol.% of the Earth.

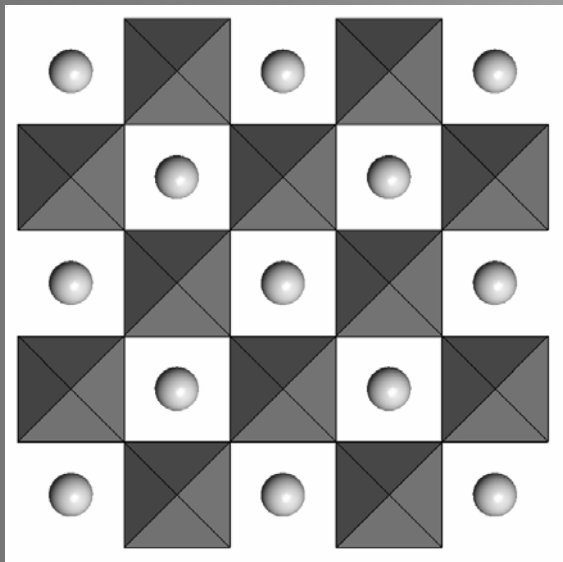
Three issues:

- ◆ **Symmetry of perovskite**
- ◆ **Decomposition of perovskite**
- ◆ **Post-perovskite phase**

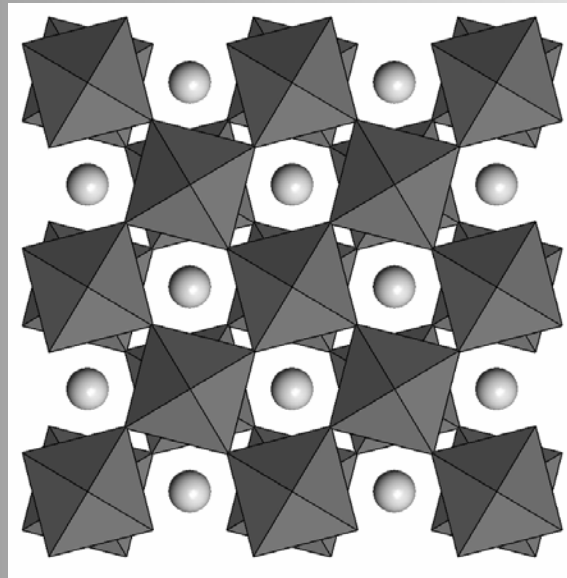
1. Symmetry



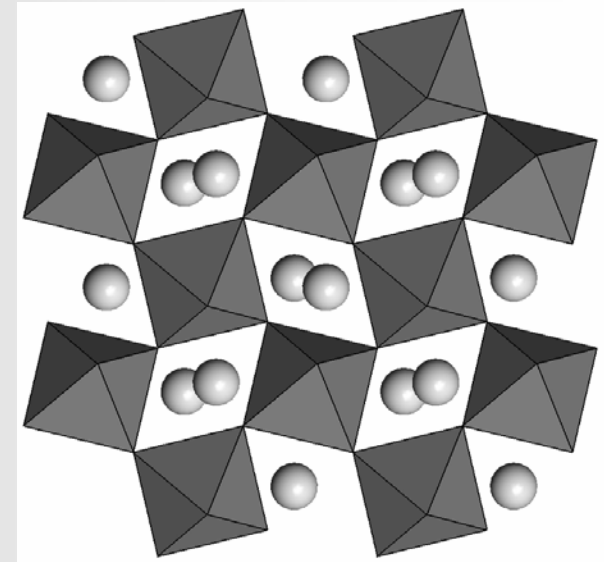
- Ideal structure – cubic ($Pm\bar{3}m$).
- $MgSiO_3$ perovskite at ambient conditions – $Pbnm$.
- Higher symmetry at high P - T ?



Cubic ($Pm\bar{3}m$)

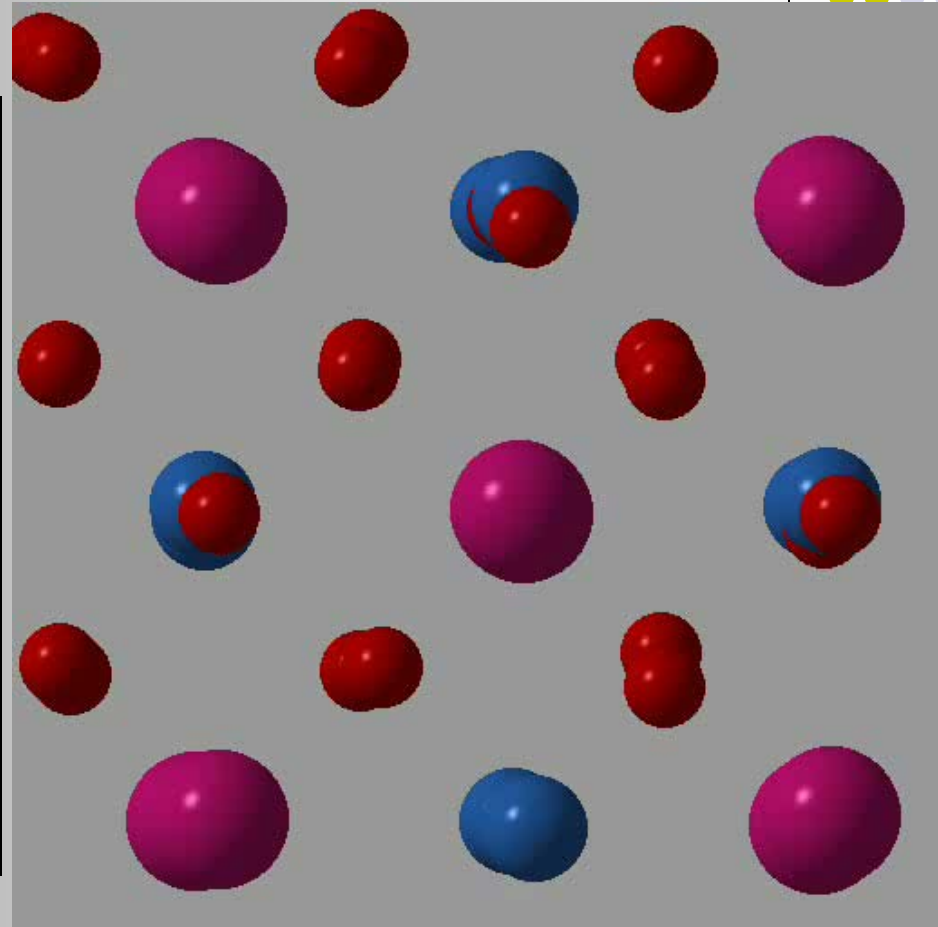
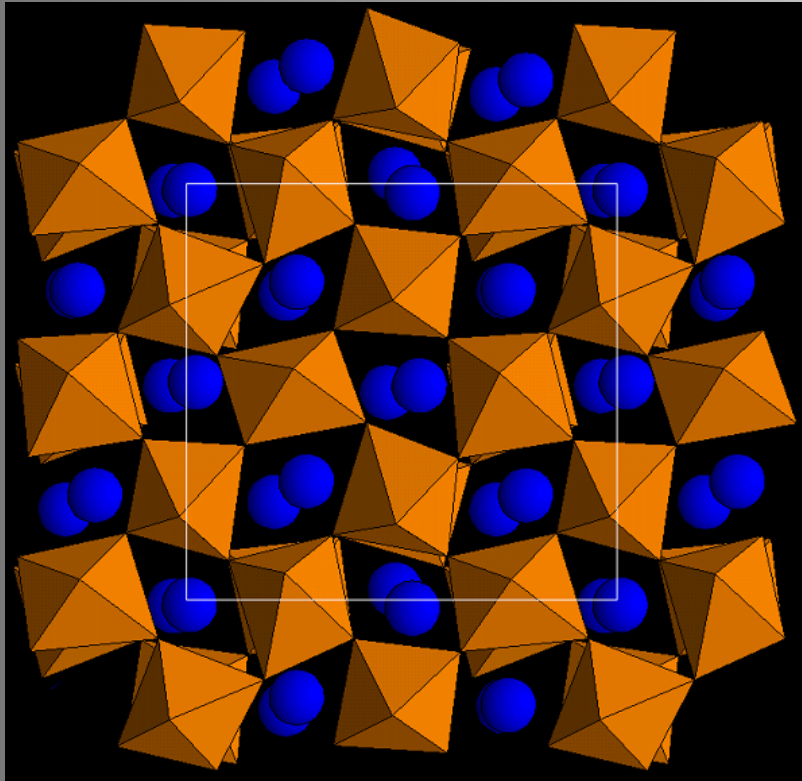
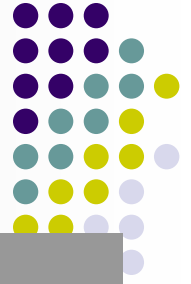


Tetragonal ($I4/mcm$)



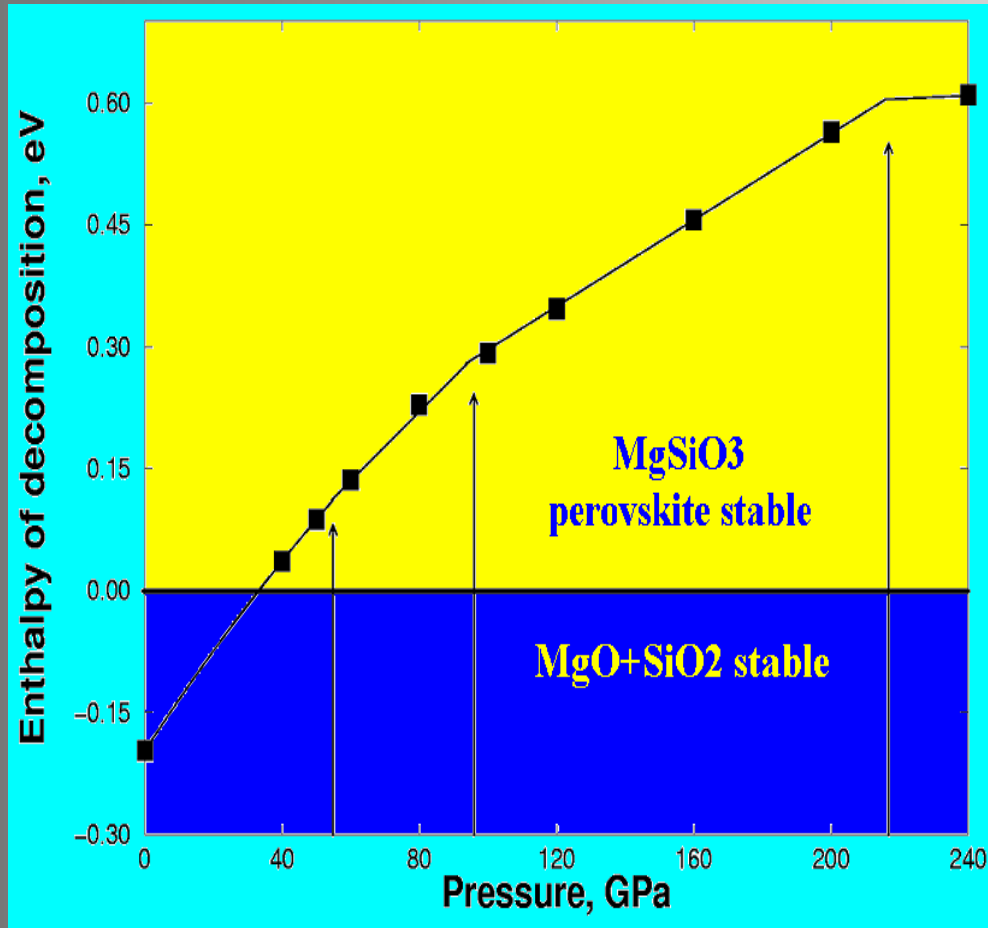
Orthorhombic ($Pbnm$)

1. Symmetry



MgSiO₃ perovskite at 88 GPa and 3500 K

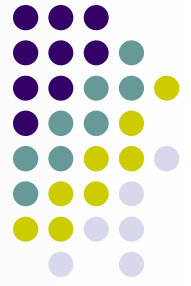
MgSiO₃: decomposition



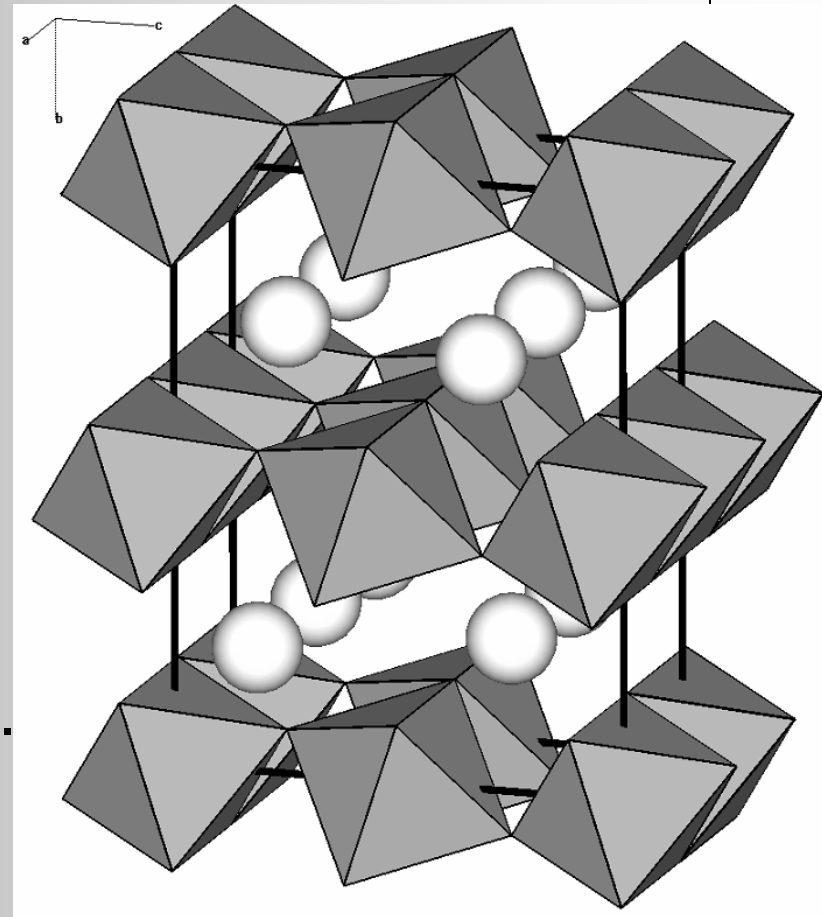
- Decomposition at high- P - T ?
- Theory: NO!
- $\Delta S = -5 \text{ Jmol}^{-1}\text{K}^{-1}$ (100 GPa, 3000 K).

Enthalpy of decomposition of MgSiO₃ perovskite. Oganov et al., In prep. (2004)

MgSiO₃: Post-perovskite phase

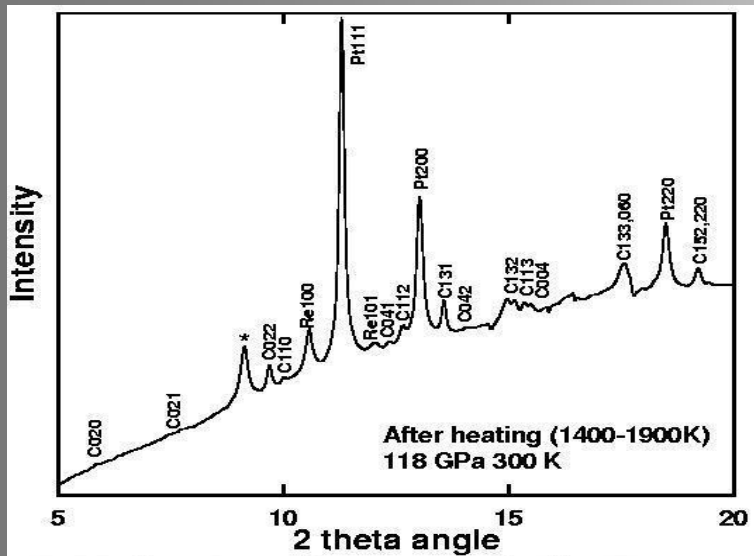
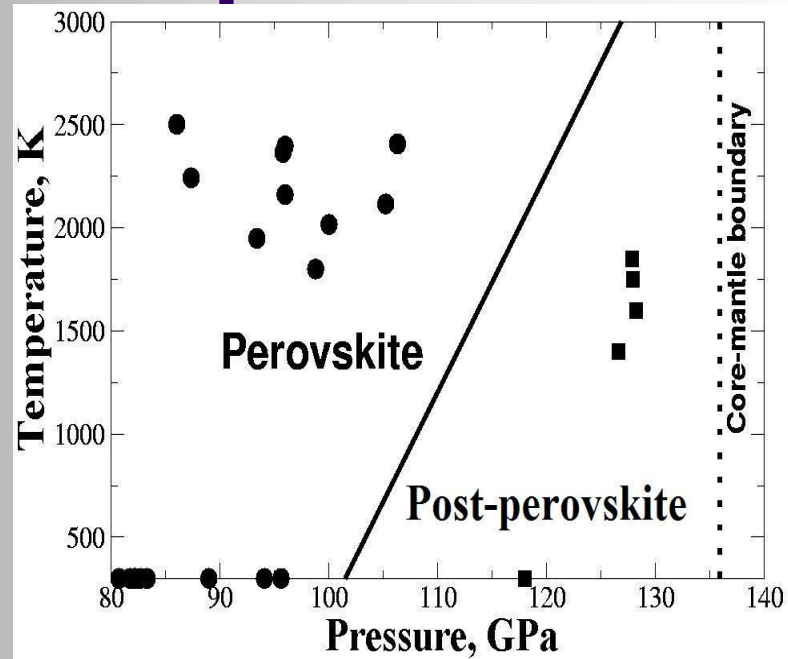
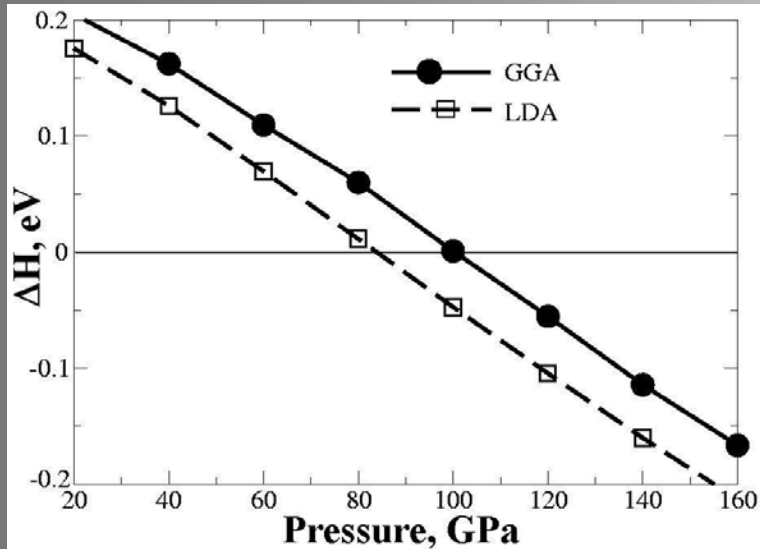


- S.Ono (2004): proposed that MgSiO₃ could adopt Fe₂O₃-III structure.
- Post-perovskite phase ~75 vol.% of D'' layer!
- Experimentally confirmed by Oganov & Ono (2004) and Murakami et al. (2004).



Structure of post-perovskite phase of MgSiO₃.
(Oganov & Ono, subm. to Nature, 2004).

Predictions and Experiment



Exp.:	PAW-GGA:
$a = 2.471(4) \text{ \AA}$	$a = 2.474 \text{ \AA}$
$b = 8.091(12) \text{ \AA}$	$b = 8.121 \text{ \AA}$
$c = 6.110(11) \text{ \AA}$	$c = 6.138 \text{ \AA}$
Space group <i>Cmcm</i>	

Elastic Constants of Post-perovskite



Table 3. Elastic constants of perovskite and post-perovskite at 120 GPa*.

	C_{11}	C_{22}	C_{33}	C_{12}	C_{13}	C_{23}	C_{44}	C_{55}	C_{66}	K	G
Perovskite	907	1157	1104	513	406	431	364	271	333	648.0	310.9
	Acoustic velocities: $v_p=14118$, $v_s=7636$, $v_\phi=11026$ m/s										
Post-perovskite	1252	929	1233	414	325	478	277	266	408	647.2	327.5
	Acoustic velocities: $v_p=14158$, $v_s=7783$, $v_\phi=10940$ m/s										

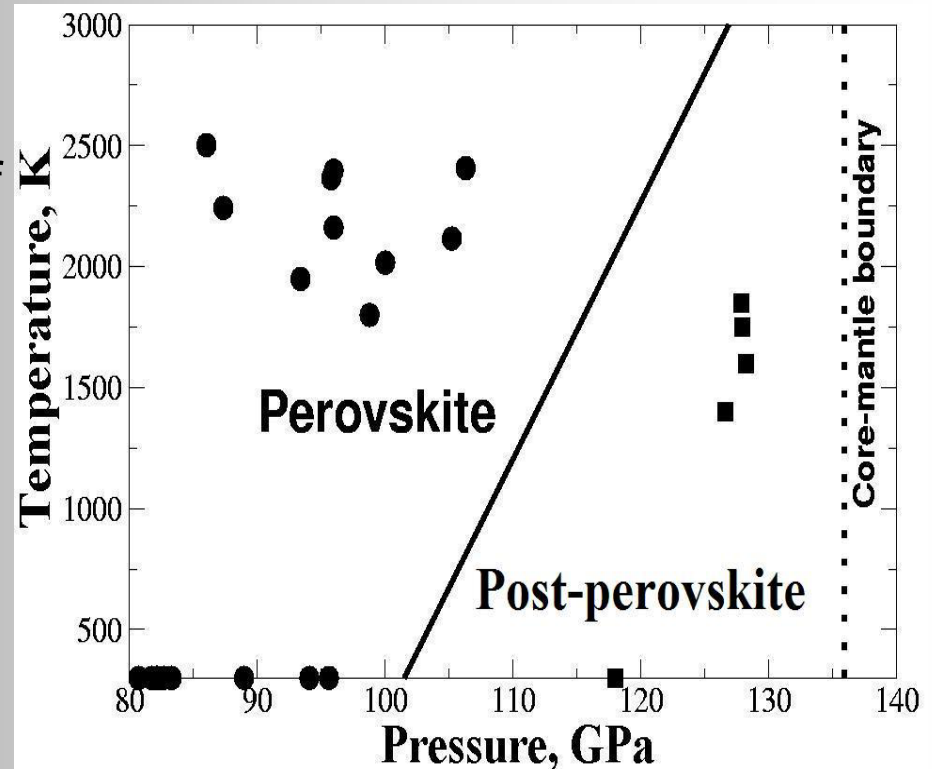
*GGA results. All elastic constants are in GPa.

- VASP: PAW-GGA calculations using stress-strain relations.
- Similar to ABINIT result using D.R. Hamann's method (when stress state is taken into account).
- Explain most of the D'' mysteries!



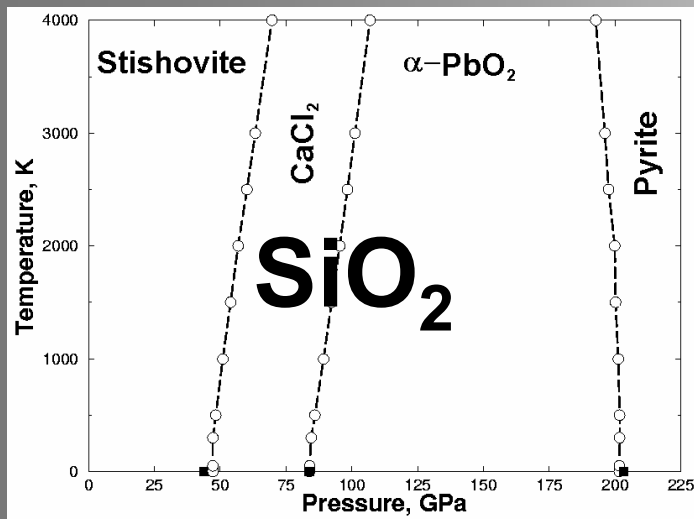
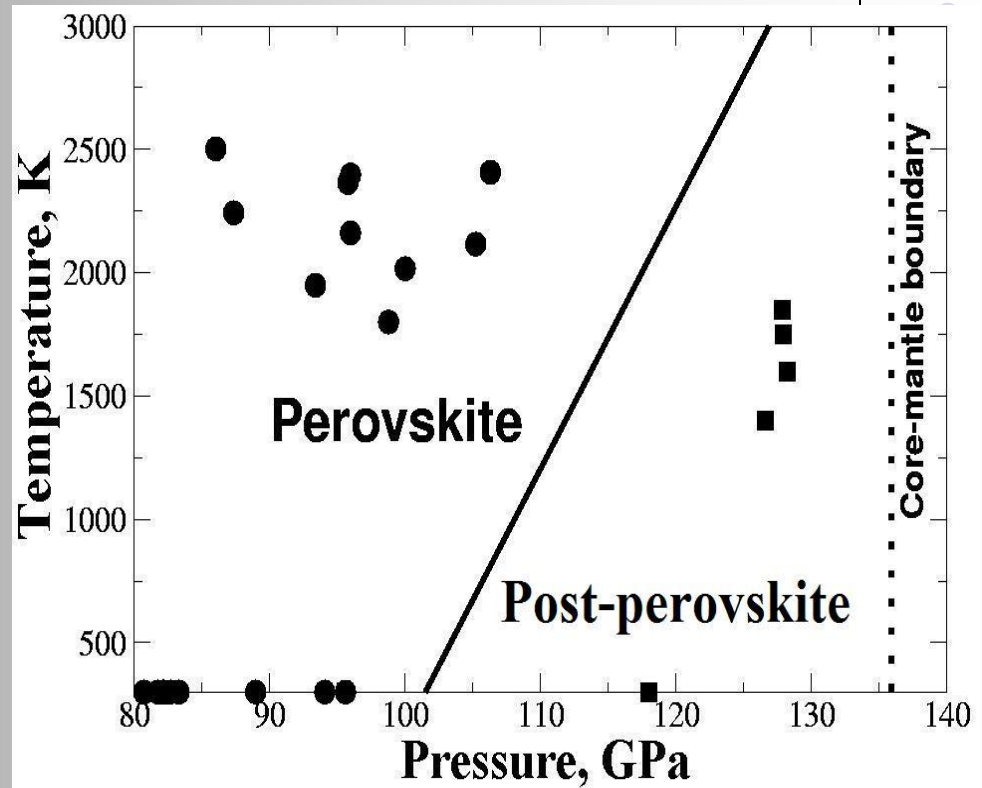
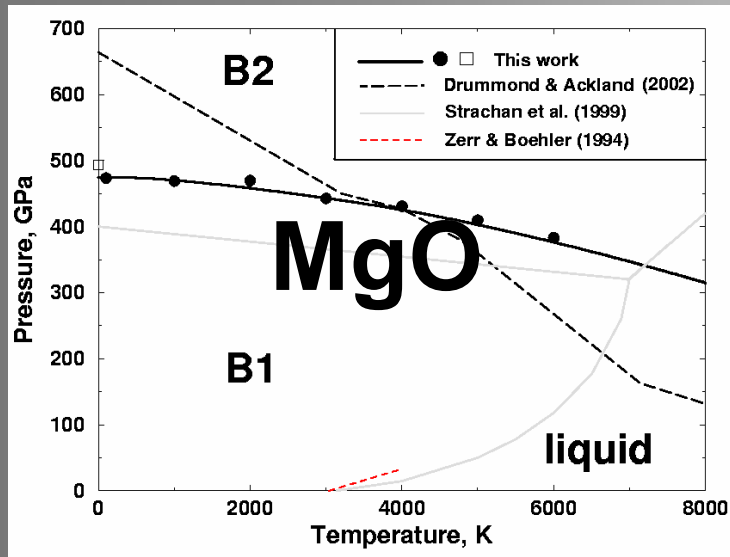
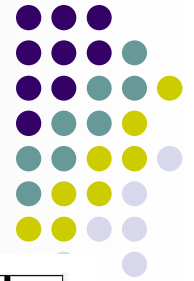
MgSiO₃: Post-perovskite phase

- Matches D'' discontinuity: 2740 km depth, 1.4% jump in v_s , strong topography.
- Explains seismic anisotropy of D'' and its other enigmas.
- D'' evolves with time.



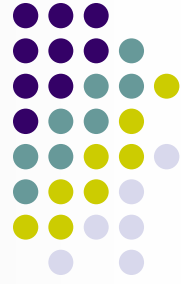
Theoretical and experimental phase diagram of MgSiO₃.

Summary



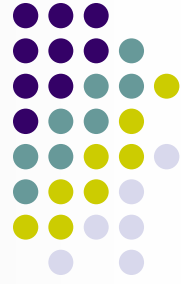
MgSiO₃

Further Work and Challenges



- Structure prediction – Genetic Algorithms?
- Solid solutions (Monte Carlo?).
- Strongly correlated systems – (Mg,Fe)O etc.



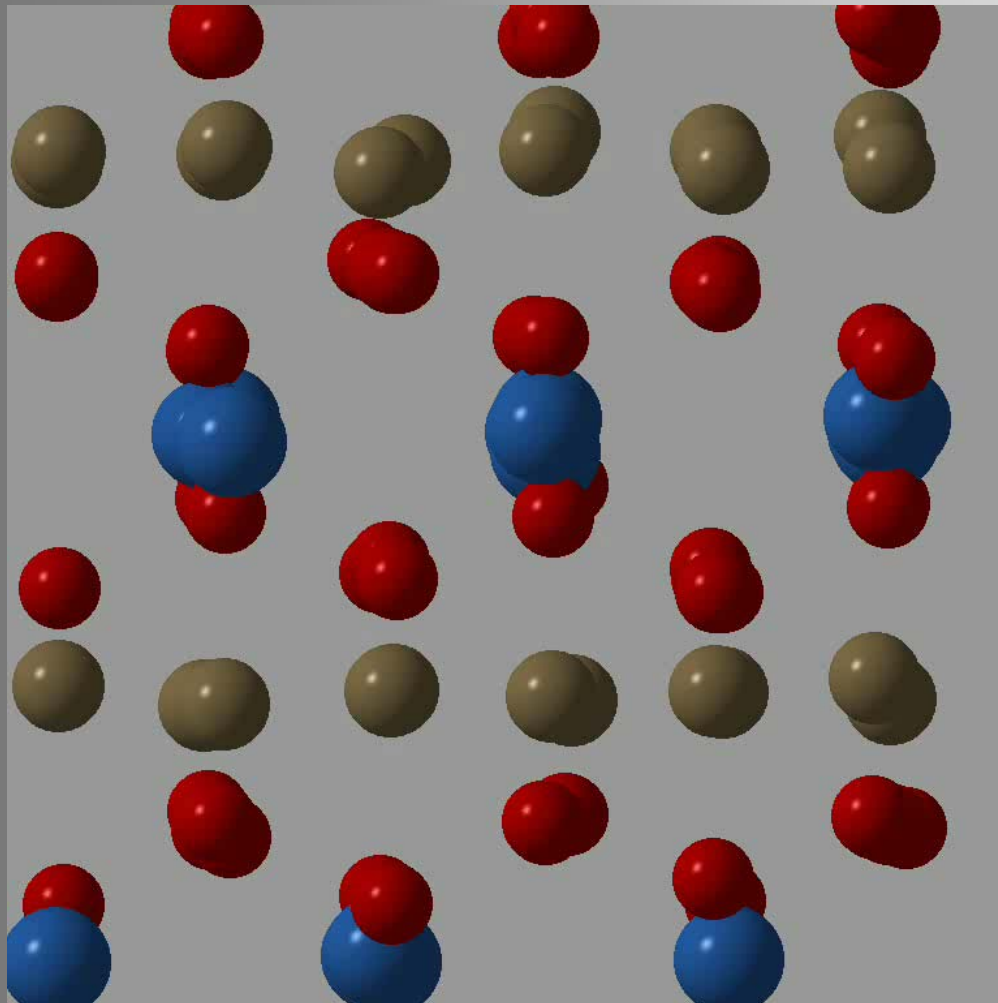
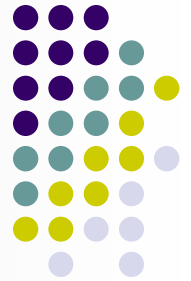


Appendix I. Accuracy.

- **MgO, 1 atm, 300 K:**
 - $C_v=36.58/36.87$ theory/exp.
 - $S = 26.81/27.13$

- **Stishovite (SiO_2), 1 atm, 300 K:**
 - $C_v=41.3/42.2$ theory/exp.
 - $S = 24.6/25.9$

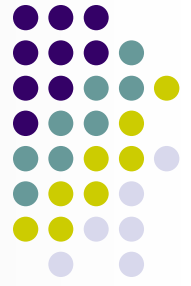
Appendix II. MD & visualisation



- 120-atom cell.
- PAW, GGA.
- VASP.

- Movie – STM3 (thanks to M.Valle & J.Favre)
- Compatible with CPMD, VASP and DL_POLY formats.
- ABINIT

Appendix III. Elasticity under pressure



- Non-uniqueness of the definition (strain type, differentiation details) when pressure (P) is present.

Most useful definition of elastic constants:

$$\sigma_{ij} = C_{ijkl} \eta_{kl}, \quad \{1\}$$

Usual definition of strain: $a_{ij}' = (\delta_{ij} + \eta_{ij}) a_{ij}^0$ {2}

$$C_{ijkl}^S = \frac{1}{V} \left(\frac{\partial^2 H}{\partial \eta_{ij} \partial \eta_{kl}} \right)_S = \frac{1}{V} \left(\frac{\partial^2 E}{\partial \eta_{ij} \partial \eta_{kl}} \right)_S + \frac{P}{2} (2\delta_{ij} \delta_{kl} - \delta_{il} \delta_{jk} - \delta_{jl} \delta_{ik}) \quad \{3\}$$

ABINIT calculates $\frac{1}{V} \frac{\partial}{\partial \eta_{11}} (V \sigma_{11})$

Example: $C_{11} = \frac{1}{V} \left(\frac{\partial^2 E}{\partial \eta_{11} \partial \eta_{11}} \right)_S = \frac{1}{V} \frac{\partial}{\partial \eta_{11}} (V \sigma_{11}) + P$ {4}